

Seismic quiescence in a frictional earthquake model

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SUMMARY

We investigate the origin of seismic quiescence with a generalized version of the Burridge–Knopoff model for earthquakes and show that it can be generated by a multi-peaked probability distribution of the thresholds at which contacts break. Such a distribution is not assumed *a priori* but naturally results from the aging of the contacts. We show that the model can exhibit quiescence as well as enhanced foreshock activity, depending on the value of some parameters. This provides a generic understanding for seismic quiescence, which encompasses earlier specific explanations and could provide a pathway for a classification of faults.

Key words: earthquake modeling, quiescence

1 INTRODUCTION

While two main seismological laws, the Gutenberg–Richter (GR) law for the magnitude distribution of earthquakes (Gutenberg & Richter, 1954) and the Omori law for the time evolution of the frequency of aftershocks (Omori, 1894; Utsu et al., 1995), are well established empirically and even may be justified theoretically (Burridge & Knopoff, 1967; Olami et al., 1992; Socolar et al., 1993; Grassberger, 1994; Rundle & Klein, 1995; Rice & Ben-Zion, 1996; Stein, 1999; Hainzl et al., 2000; Helmstetter et al., 2004; Hergarten & Krenn, 2011; Serino et al., 2011; Jagla, 2010; Jagla, 2010b; Jagla, 2013; Braun & Peyrard, 2013; Mega et al., 2003; Kawamura et al., 2012; Pelletier, 2000), our knowledge about foreshocks is still limited. If a clear pattern could be distinguished in the foreshock activity it might provide an early warning for earthquakes, but we are far from this stage because the observations reveal very different pictures. Many earthquakes are preceded by foreshocks, but their frequency can nevertheless vary widely depending on the type of earthquake (Bouchon et al., 2013). Moreover, some earthquakes are preceded by an unexpected calm period, lasting for several hours or more (Jones & Molnar, 1979). It is such a period of quiescence, viewed as a characteristic feature of the imminence of the main shock that allowed the only successful prediction of an earthquake (Raleigh et al., 1977), which saved a large number of lives in China in 1975.

Although the mechanisms that generate foreshocks are still under discussion (Bouchon et al., 2013), the existence of some activity before a major earthquake (MEQ) does not seem surprising. In recent experiments with slider pushed on the side, Fineberg *et al.* (Rubinstein et al., 2004; Rubinstein et al., 2007; Rubinstein et al., 2011) showed that the global sliding, which corresponds to a MEQ, is preceded by propagation of several cracks called precur-

sors that may be used as indicators of the upcoming MEQ. This is also consistent with some ideas that relate earthquakes to self organized criticality or cascade mechanisms (Olami et al., 1992; Socolar et al., 1993; Grassberger, 1994; Helmstetter et al., 2004; Hergarten & Krenn, 2011).

The fact that a quiescence period could be a characteristic announcement of a MEQ looks more surprising. Various mechanisms have been considered to explain seismic quiescence (Hergarten & Krenn, 2011; Main & Meredith, 1991). Here we propose a generic mechanism related to the distribution of the threshold for the breaking of the contacts along a fault and illustrate it by simulations of a simple earthquake model.

The Burridge and Knopoff (BK) spring-block model (Burridge & Knopoff, 1967) further developed in a number of works (Olami et al., 1992; Hainzl et al., 2000; Helmstetter et al., 2004; Hergarten & Krenn, 2011; Serino et al., 2011; Jagla, 2010; Jagla, 2010b; Jagla, 2013; Braun & Peyrard, 2013; Braun & Tosatti, 2014; Kazemian et al., 2015) (see also reviews (Pelletier, 2000; Kawamura et al., 2012) and references therein) has been used as a generic model which reproduces many features of earthquakes. Simulations showed that some generalized versions of the BK model may demonstrate the GR law and even the Omori law, but the existence of foreshocks in this model is not clearly demonstrated and explained yet (e.g., see (Hainzl et al., 2000; Helmstetter et al., 2004; Hergarten & Krenn, 2011; Jagla, 2010; Jagla, 2010b; Jagla, 2013; Kazemian et al., 2015; Pelletier, 2000)).

In BK-type models the top block (the slider) is coupled with the bottom block (the base assumed to be fixed) by a set of frictional contacts. When the slider moves, a frictional contact i , modeled as an elastic spring, elongates so that the force at the contact point, $f_i = kx_i$ (we assume that all contacts have the same rigidity k for the sake of simplicity), increases until it reaches a threshold value f_{si} for which the contact breaks. Then it forms again, with zero stretching $x_i \sim 0$, and the process can repeat. The evolution

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of such a system is determined by the distribution of thresholds $P_c(f_{si})$. The general features of this distribution also determines the nature of the foreshock activity, as one can realize by considering two simple examples.

i) If $P_c(f_{si})$ is one-peaked (Braun & Peyrard, 2010), e.g., Gaussian with the center at f_s and a standard deviation Δf_s , then the total frictional force on N contacts, $F = \sum_{i=1}^N f_i$ linearly increases with the slider displacement X until the contact forces reach values $f_i \sim f_s - \Delta f_s$. Then some contacts start to break, one by one, giving rise to small earthquakes. The growth of $F(X)$ slows down and turns into a decay for larger values of X , and under certain conditions (Braun, 2015) an elastic instability may occur. The top block slides for a distance $\Delta X \gtrsim \Delta f_s/k$ while most of the contacts break. This event corresponds to a large earthquake. In this case the earthquake activity increases just before the MEQ. Such a scenario is common in natural earthquakes (Kazemian et al., 2015; Bowman & King, 2001; Jordan & Jones, 2010).

ii) If the threshold distribution is two-peaked, e.g., when the interface consists of two types of contacts, weak and strong ones with thresholds f_{s1} and f_{s2} respectively, as recently considered in (Kazemian et al., 2015), the scenario is qualitatively different. To see how it works let us consider a simple case. We assume that, when a contact breaks it forms again as a weak contact with probability p and strong contact with probability $1-p$, and that the breaking of a contact reduces the friction force by φ . At time $t_0 = 0$ we start with $n_1^{(0)} = pN$ weak contacts and $n_2^{(0)} = (1-p)N$ strong contacts. If the slider moves at velocity v the driving force grows as $Nkvt$. When it reaches f_{s1} at time t_1 , all weak contacts break. The friction force drops by $\varphi n_1^{(0)}$ and the broken contacts reform, leading to $n_1^{(1)} = pn_1^{(0)} = p^2N$ weak contacts and a growing number of strong contacts. The same process can repeat again, leading to an infinite series of foreshocks, while the number of weak contacts decreases as $n_1^{(j)} = p^{j+1}N$. During the time interval $t_j - t_{j-1}$ the force has to grow enough to reach f_{s1} again after the breaking $n_1^{(j-1)}$ weak contacts. It requires a time $t_j - t_{j-1} = \varphi p^j N / (Nkv)$ so that the total time for the infinite series of foreshocks stays finite $\sum_{j=1}^{\infty} (t_j - t_{j-1}) = (\varphi/kv)p/(1-p)$. After that time only strong contacts persist. As the slider keeps moving, there is a calm period without any foreshock until the strong threshold f_{s2} is reached. Then all strong contacts break suddenly, leading to a MEQ.

This picture of contacts with a two-peaked distribution is oversimplified, and would not show all features of real earthquakes, such as the GR law or the Omori law for aftershocks, but it shows how the properties of the contact thresholds are sufficient to lead to seismic quiescence. Let us now consider a more realistic case and investigate its properties.

2 MODEL

We consider the earthquake model previously developed in Braun & Tosatti (2014) and Braun & Scheibert (2014). It is based on the BK-type model, i.e. a set of frictional contacts between a slider and a base, but with the following two important ingredients.

First, by contact we mean a macro-contact consisting of a large number of micro-contacts on the area λ_c^2 , where $\lambda_c \approx a^2 E/k$ is the elastic correlation length (Caroli & Nozieres, 1998; Braun et al., 2012) (here a is the average distance between the micro-contacts and E is the Young modulus of the slider). For a frictional metal/metal interface λ_c is typically of the order of μm (Caroli & Nozieres, 1998; Braun et al., 2012), but for a seismic fault it may

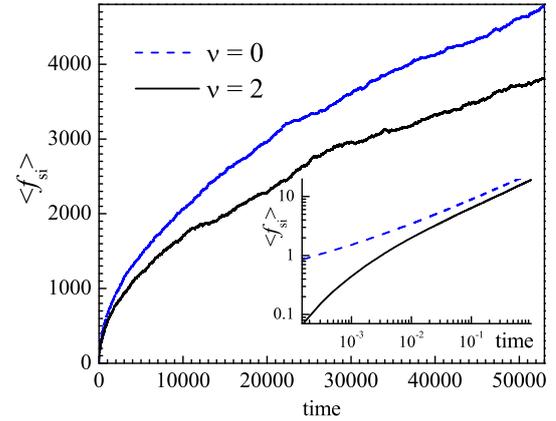


Figure 1. (Color online) Aging of macro-contacts: growth of the average contact threshold $\langle f_{si} \rangle$ with its lifetime for $\beta = 100$, $\varepsilon = 2/\Delta f_s^2$, $\nu = 0$ (blue dashed) and $\nu = 2$ (solid curve), $f_s = 1$, $\Delta f_s = 0.3$, $N = 3001$. The inset shows the short-time behavior in log-log scale ($N = 100001$).

be much larger. On the distance λ_c the slider may be treated as rigid, while at larger distances we have to account for the deformation of the block. Thus, in our model a contact is a macroscopic object. It has its own “degrees of freedom”, which may lead to its aging resulting in a growth of its threshold value as the lifetime of the contact increases. Contact aging is a complex stochastic process which can have various physical or chemical origins. We model it by a simple Langevin equation (Braun & Peyrard, 2013; Braun & Tosatti, 2014)

$$df_{si}(t)/dt = B(f_{si}) + G\xi(t), \quad (1)$$

where $B(f)$ and G are the so-called drift and stochastic forces respectively (Gardiner, 1985), and $\xi(t)$ is a Gaussian random force. If we chose

$$B(f) = \left(\frac{2\pi f_s}{t_0} \right) \beta^2 \frac{(1 - f/f_s)(f/f_s)^\nu}{1 + \varepsilon(f/f_s)^{\nu+2}}, \quad (2)$$

and

$$G = (4\pi/t_0)^{1/2} \beta \Delta f_s, \quad (3)$$

where β , ε and ν are dimensionless parameters, the threshold values grow with time as shown in Fig. 1. For $\varepsilon = 0$ and $\nu = 0$ the stationary solution of the Langevin equation leads to a distribution of thresholds $P_c(f_{si})$ which is a Gaussian centered at f_s with half-width Δf_s , while for $\varepsilon > 0$ the stationary distribution of thresholds has a power-law tail. The factor $(f/f_s)^\nu$ introduces a “delay” in contact formation as demonstrated in Fig. 1 (inset). A newborn contact is very weak and initially its threshold grows faster than in the asymptotic limit.

The choice $\varepsilon(\Delta f_s/f_s)^2 \sim 1$ leads to the GR law with the exponent $b \sim 1$ (Braun & Peyrard, 2013). Note that an aging of the thresholds is a necessary condition for the existence of a stick-slip behavior of a sliding contact (Braun, 2015).

For spring-block models, the hypothesis made on the properties of the contacts are crucial to determine the behavior of the model. While few studies were devoted to foreshocks, many model developments were derived from the goal of reaching a proper description of aftershocks. It was soon realized that contact aging is important. Since the pioneering work of (Dieterich, 1972) this is one of the most accepted mechanisms of aftershocks. The first approach was to explicitly introduce a time-dependent friction co-

efficient using an expression deduced from experimental observations (Dieterich, 1979a; Dieterich, 1979b). This approach further evolved in a rate- and state-dependent representation of the fault constitutive properties (Ruina, 1983; Marone, 1988; Dieterich, 1994). In these approaches the friction coefficient depends on a state variable θ and the model is completed by a deterministic differential equation which determines the time dependence of the state variable. The meaning of this state variable depends on the model. It can be the time of contact or velocity (Marone, 1991; Dieterich, 1994), but it can also be the local slip (Ruina, 1983). Recent studies testing these two approaches (Helmstetter & Shaw, 2009) have shown that both evolution laws produce qualitatively similar behaviors (afterslips, slow earthquakes, aftershocks) but that the slip law is more unstable than the aging law. Velocity strengthening has also been involved to explain some features of post-seismic slip (Perfettini & Ampuero, 2008). Introducing a state variable appeared therefore as an essential step to describe the experimental observations on earthquakes, in particular the aftershocks and their distribution. However the diversity of the proposals for the evolution of the state variable shows that the choice of the best description is still an open problem. Our approach proceeds along the same idea that a contact has internal degrees of freedom, with two major differences (Braun & Peyrard, 2013). First the internal state of a contact evolves according to a stochastic equation, Eq. (1), instead of a deterministic differential equation. This takes into account the influence of external phenomena, such as for instance some tremors in the earth crust that could come from far away, and it models the fluctuations at the scale of the local contacts. Second, the description of a contact is based from a underlying physical model at a smaller scale. The idea is that, as explained above, at the scale of a fault a contact is actually a macro-contact which is the result of a large number of local contacts. Within this viewpoint, the laws of friction can be established from a master equation describing the breaking and re-forming of the many micro-contacts (Braun & Peyrard, 2010). The interest of this approach that it provides a basis to include the physical properties of the local contacts, for instance the formation of chemical bonds, or local plasticity, in the properties of the macro-contacts which enter in the spring-block model, instead of postulating an equation for the state variable.

Second, we incorporate the elasticity of the slider. The standard BK model assumes that, when a contact breaks, the released stress is arbitrarily redistributed among neighboring contacts. Actually this redistribution is due to the elastic interaction between the contacts and depends on the three-dimensional deformation of the slider, which cannot be rigorously treated, neither analytically nor numerically, unless the full geometry of the fault is taken into account. Instead, here we use the one-dimensional model, where the slider is divided in two layers as proposed in (Braun & Scheibert, 2014; Braun & Tosatti, 2014), which gives a reasonable approximation of the exact behavior.

In this model shown schematically in Fig. 2, the bottom layer is divided into rigid λ_c^3 -cubes coupled together by springs of rigidity $K = E\lambda_c$ and coupled by frictional springs of rigidity k with the base. It describes the “interface” layer (IL). The other part of the slider, the upper layer (UL), is divided into a chain of parallelepipeds of height $N_L\lambda_c$ coupled by springs $K_L = N_LE\lambda_c$. The UL and IL are coupled by the set of N transverse springs $K_T = E\lambda_c/[2(1 + \sigma_P)N_L]$, where σ_P is the slider Poisson ratio. Note that the geometry of the blocks is introduced to allow us to make a link between the elastic constants of the model (K , K_T and K_L) and the elastic parameters of the sliding medium. It

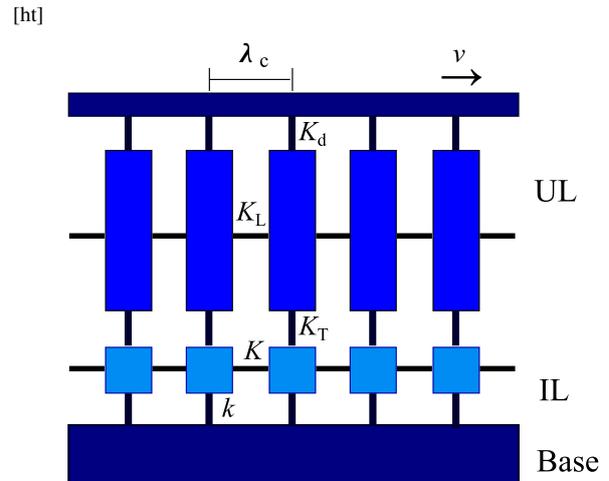


Figure 2. The model. The upper layer (UL) is split in rigid blocks of size $\lambda_c \times \lambda_c \times N_L\lambda_c$ connected by springs of elastic constant K_L . The interface layer (IL) is split in rigid blocks of size $\lambda_c \times \lambda_c \times \lambda_c$ connected by springs of elastic constant K . The UL and IL are coupled by springs of elastic constant K_T . The IL is connected with the rigid bottom block (the base) by contacts, represented by “frictional” springs of elastic constant k , which break when the local stress exceeds a threshold value. The UL is driven with the velocity v through springs of elastic constant K_d .

is not an essential feature of the model. In this model, if we push the most-left block of the UL until the most-left frictional contact breaks, a crack emerges and propagates through the interface for a finite distance $\Lambda \gg \lambda_c$ till it is arrested in qualitative agreement with experiments (Rubinstein et al., 2004; Rubinstein et al., 2007; Rubinstein et al., 2011). The UL plays the role of a reservoir, where the elastic energy is stored and partially released at the main shock, while the unreleased part of the elastic energy results in aftershocks satisfying the Omori law (Braun & Tosatti, 2014).

Finally, we attach springs K_d to the top boundary of the UL and drive the ends of these springs with a velocity v to simulate the driving of the seismic fault. For the simulations we use periodic boundary conditions in the direction of the chain of blocks.

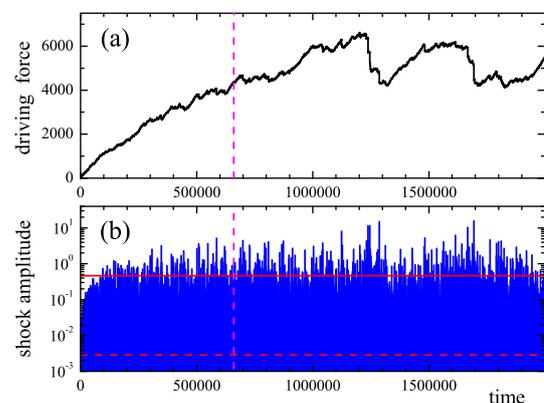


Figure 3. (Color online) Time evolution of the system: (a) the frictional force $F(t)$, and (b) the (global) amplitude of EQs $\mathcal{A}(t)$ versus time.

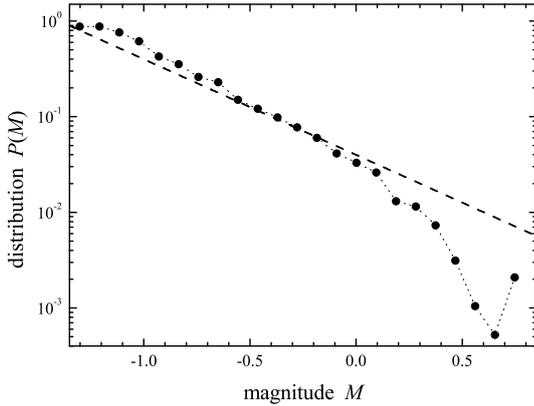


Figure 4. Statistics of EQ magnitudes presented in Fig. 3 showing the Gutenberg-Richter power law behavior; the dashed line corresponds to the exponent $b = 1$.

3 SIMULATION RESULTS

The simulation results are presented in Figs. 3–6. For the chosen set of parameters, we ran 20 independent calculations using different sets of random numbers, and with the protocol described in the Material and method section we obtained 10.8×10^6 shocks above the background level and analyzed 7012 main shocks which occurred during the time interval $T = 0.67 t_{\max} \approx 1.34 \times 10^6$ after the first 33% of data were discarded in each simulation.

The dependencies of the driving force $F(t)$ (Fig. 3a) and the amplitude of shocks $\mathcal{A}(t)$ (Fig. 3b) on a large time scale both look rather stochastic. As we observed with shock visualization during simulation, and as can also be guessed from comparison of Figs. 1 and 3a, the frictional force is mainly determined by one or few large thresholds with $f_{si} \gg f_s$ that survived for a long enough time; the breaking of these contacts results in MEQs.

The statistics of the shock amplitudes follows the GR power law (see Fig. 4) for an interval of magnitudes $\Delta M \gtrsim 2$, which is however limited by largest possible magnitude determined by the ratio of the aging rate β and the driving velocity v , $\mathcal{M}_{\max} \propto \log_{10}(\beta/\sqrt{v})$ (Braun & Peyrard, 2013) (as we use here a rather small quasi-1D system, we cannot expect a very large value of ΔM).

Figure 5 shows the $\mathcal{A}(t)$ dependence for the last 5% of the simulation time (top panel) and more detailed pictures at a finer time scale of two typical MEQs (middle panel); the bottom panel of Fig. 5 shows the color maps of the earthquake amplitude.

Figures 6 presents the main result of the data analysis, the average number of shocks $n(\tau)$ in an interval $\delta\tau$ ($\delta\tau \approx 0.0648$), versus τ the rescaled time interval from ($\tau < 0$), or after ($\tau > 0$), the corresponding main shock. We show both the mean value $n(\tau)$, computed for the 7012 analyzed shocks, and its fluctuations for different MEQs, measured by its standard deviation over all analyzed MEQs. The variation of the fluctuations versus τ are the most interesting. Before the MEQ they are large, and vary widely with τ in the range $\tau \lesssim -0.7$, showing a significant and random foreshock activity. Then suddenly they drop to a small value and stay so for the last period before the MEQ, $-0.7 \lesssim \tau < 0$. In this range none of the 7012 analyzed MEQs shows any noticeable foreshock activity. As expected, after the MEQ, Fig. 6 shows a large aftershock activity.

Figure 7 shows that the sharp drop of foreshock activity before large events is clearly due to the aging of the contacts, de-

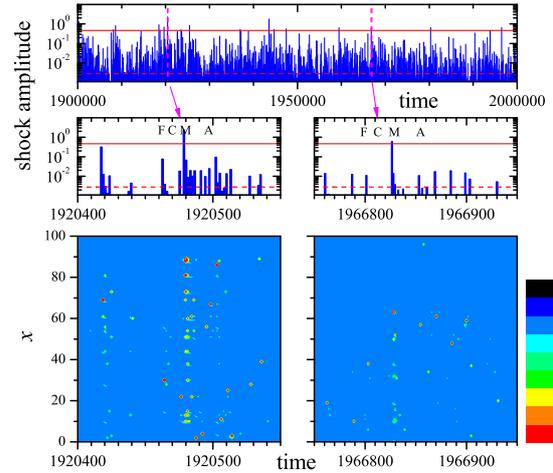


Figure 5. (Color online) Typical earthquakes in the model. Top panel: the EQ amplitude $\mathcal{A}(t)$ versus time for the last 5% of simulation time. Middle panel: detailed $\mathcal{A}(t)$ dependencies for two time intervals. Bottom panel: the corresponding color maps of the EQ amplitude on the (t, x) plane, using the color scale indicated on the right. F—foreshocks, C—calm, M—main shock, A—aftershocks.

scribed by the parameter β in Eq. (2). When β is reduced from $\beta = 100$ to $\beta = 3$ ($\beta = 0$ corresponding to no aging at all), all other parameters being preserved, the calm period disappears and a slight increase of activity immediately before the MEQ is even observed. We can also notice that, in the absence of aging, the fluctuations of the foreshock and aftershock activity with time, are much smaller. Similarly the standard deviations of $n(\tau)$ between the different main shocks are smaller and almost time-independent. The joint evolution of the aftershocks and foreshocks when β varies shows some similarity with the correlation between the rates of foreshocks and aftershocks found in earthquakes catalogs in different geographic regions (Lippiello et al., 2017).

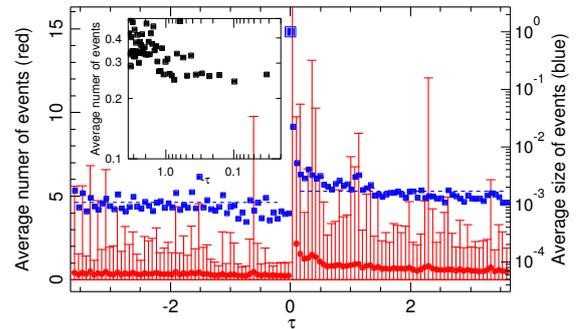


Figure 6. (Color online) Foreshocks and aftershocks statistics in the presence of contact aging ($\beta = 100$): the rate of fore- and aftershocks $n(\tau)$ (red solid symbols) and their associated standard deviations (red bars), and the shock amplitudes $\mathcal{A}(\tau)$ (blue squares) relative to the corresponding main shock. The horizontal dash lines show the average magnitudes of the foreshocks and aftershocks. The inset shows the average number of foreshocks $n(\tau)$ using a log-log magnified scale. The results plotted on this figure have been obtained from the analysis of 7012 main shocks, collected over 20 independent simulations.

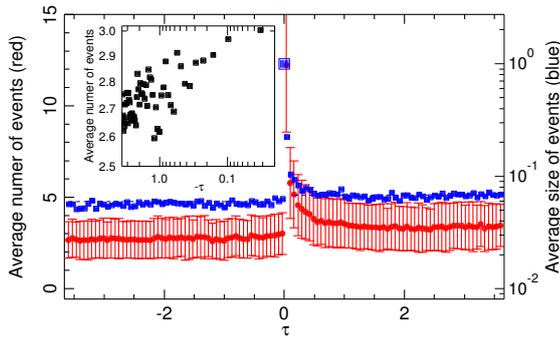


Figure 7. (Color online) Same as Fig. 6 with $\beta = 3$, i.e. with very little aging of the contacts. In this case the magnitudes of the main events are not as large as in the presence of aging, so that we define the main shocks by $\mathcal{A}_{\text{MEQ}} = \kappa \mathcal{A}_{\text{max}}$, with $\kappa = 0.4$. A set of 1205 MEQs was used to draw this figure. The inset shows the average number of foreshocks $n(\tau)$ using a log-log magnified scale.

Relation with the actual scales of earthquakes. To couple our dimensionless units with real ones, we use the following estimation (Helmstetter et al., 2004): let us calculate the average time between the MEQs as $t_{\text{rec}} = T/n_{\text{MEQ}} \approx 2.8 \times 10^3$. Then, if we identify this time span with, e.g., 100 days (which is the recurrence time of $\mathcal{M} > \mathcal{M}_0 = 5$ earthquakes in Southern California), we obtain the correspondence between time units in our model and in real seismicity:

$$t_0 = 1 \longleftrightarrow \tau_0 \equiv t_0/\alpha \approx 3.8 \times 10^{-3} \longleftrightarrow t_{\text{real}} \sim 1 \text{ hour.}$$

Thus, the time span in Fig. 3 corresponds to about 230 years and the average calm period before the MEQs, $\tau_{\text{calm}} \lesssim 0.7$ observed in Fig. 6, corresponds to about 7 days. However, this estimation depends on the value of \mathcal{M}_0 we choose. For example, at the Parkfield segment along the San Andreas fault, California, interplate earthquakes of magnitude about $\mathcal{M}_0 = 6$ have occurred at recurrence intervals of 23 ± 9 yr (Kawamura et al., 2012), while along the Nankai trough, where the Philippine Sea plate subducts beneath southwestern Japan, great earthquakes of magnitude $\mathcal{M}_0 = 8$ have repeatedly occurred every 100 years (Kawamura et al., 2012). If earthquakes in a given region follow the GR law, then the recurrence period depends on \mathcal{M}_0 as $t_{\text{rec}} \propto \exp(\mathcal{M}_0)$. In a simple model we cannot provide an unquestionable value of the magnitude that could be compared with actual magnitudes, but the main difficulty for a comparison with actual data is that the limited size of the model restricts the range of the observable events. Figure 6 shows events with a range that spans less than three orders of magnitude, while actual foreshocks can be much smaller with respect to a MEQ. The model cannot show the smallest events and tends to overestimate the duration of the calm period. In actual earthquakes it may sometimes only last for a few hours or less than a day as in (Raleigh et al., 1977), but may extend to much longer periods reaching several months (Main & Meredith, 1991).

4 DISCUSSION AND CONCLUSION

In order to correctly describe various features of earthquakes, the model that we used in this work is more complex than the simplified case that we presented in the introduction, and the probability

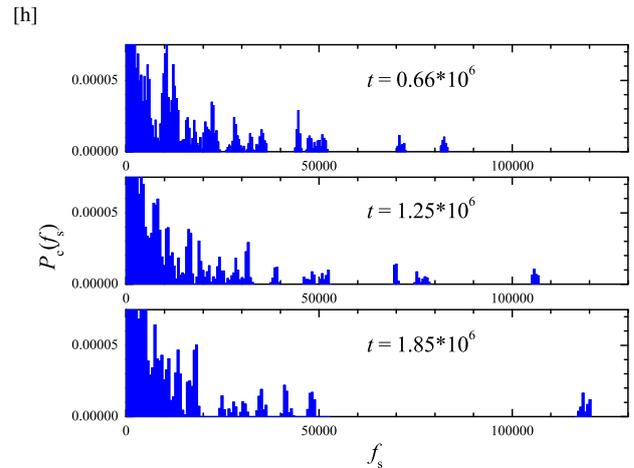


Figure 8. Evolution of the probability distribution of the contact thresholds as time grows (from top to bottom figure, as indicated in each frame).

distribution of the breaking thresholds is not imposed *a priori* but instead results from the dynamics of the model through Eq. (1). Figure 8 shows that the qualitative mechanism for the origin of a calm period that we presented in the introduction is nevertheless also present in this more elaborate case. While the initial probability distribution of the thresholds for contact breaking $P_c(f_{si})$ was a Gaussian, the evolution of the thresholds according to Eq. (1) leads to a very different shape as time evolves. The maximum value of f_{si} grows, and moreover the distribution tends to split into a large cluster of contacts which break easily (the “weak contacts” of the qualitative model) and a few very strong contacts. The contacts with intermediate thresholds tend to disappear. It is the few strong contacts that prevent the sliding and are responsible for the MEQs when they finally break. The gap between these strong contacts and the many weak contacts is responsible for the calm period. This evolution is mostly governed by the aging of the contacts.

In natural seismicity quiescence is observed in about 40% of events (Turcotte et al., 2000; Kawamura et al., 2012): the frequency of small events is gradually enhanced preceding the main shock, whereas, just before the main shock, it is suppressed in a close vicinity of the epicenter of the upcoming event. In about 60% of natural earthquakes another scenario is observed. Large earthquakes are preceded by a period during which the surrounding region experiences a phase of accelerated seismic release (ASR) (Ben-Zion & Lyakhovsky, 2002; McGuire et al., 2005). The rate of these foreshocks also follows the Omori law (Kagan & Knopoff, 1978; Jones & Molnar, 1979), usually with a smaller value of t_0 . Our model may also demonstrate such a behavior for smaller values of the ratio β^2/v (Braun & Tosatti, 2014), when the distribution of large thresholds is more uniform. However, in the model as in reality, earthquakes exhibit a broad dispersion of their features from even to event. This is why seismic quiescence is noticeable on the standard deviations of the foreshock activity much more than on the mean value of $n(\tau)$ averaged over a large number of MEQs (Inset of Fig. 6). The intensity of the foreshock activity prior to the calm period varies strongly from event to event and can even stay rather low for some MEQs. This means that using quiescence as a warning signal may be unreliable even for a fault which have the characteristics which lead to quiescence.

In nature large earthquakes are almost always followed by a very large seismicity rate (aftershocks). In our model aftershocks

are less likely for parameters that lead to a calm period before big events (see Figs. 6 and 7). This is consistent with the picture shown in Fig. 8 because the quiescence is associated to an exhaustion of the weak contacts before the MEQ so that the large event is more likely to fully release the stress in the fault. It would be interesting to check whether such an effect is actually observed for natural earthquakes showing quiescence. However the decrease of aftershock activity may be exaggerated in our calculations due to the restricted size of the model. In an actual earthquake a complete fault line does not break at once, and instead the partial breaking tends to accumulate stress in some regions, as our model does for some parameter sets.

This study points out the crucial role of the distribution of thresholds to control the pattern of foreshocks. The existence of a few strong contacts tends to induce a period of seismic quiescence before a large earthquake. Such a distribution could of course come from the geometry of a fault. Some studies correlated different foreshock patterns to general features of faults, such as interplate or intraplate earthquakes (Bouchon et al., 2013). However, one important result of our study is that the aging of the contacts can also be a determining factor, in slowly building a multi-peaked distribution of contact thresholds. Several mechanisms may be responsible for aging of the sliding interface, such as removing of dislocations from asperities, growth of contact sizes due to plastic deformation, squeezing of a “lubricant” (e.g., water or powdered rock) from the interface or its solidification due to high pressure, coalescence of contacts, etc. These features may be characteristic of a particular fault or a type of rock found in some area, so that statistics of quiescence could differ from place to place. As the time scale of quiescence is short compared to the periods generally involved in foreshock studies, there are very few available data which could support this idea. However, a recent article (Lippiello et al., 2017) presents [in Fig.3] statistics of foreshocks which show some drop of activity before a main shock for the Northern California Earthquake Catalog (RNCEC) (particularly for magnitude 2) which does not appear for other catalogs. This could suggest that such local differences exist although this is to be taken with caution because the study did not pay a particular attention to quiescence. A recent study has stressed the dominant role of the formation of interfacial chemical bonds (Qunyang Li et al., 2011). It is interesting that the chemistry is also proposed as a cause for seismic quiescence (Main & Meredith, 1991). It would be interesting if geologists could correlate the chemical properties of the rocks involved in the faults and the pattern of foreshocks.

Real earthquakes, as well as the model that we have presented when its parameters are changed, can exhibit different patterns of foreshocks. An acceleration of the events prior to a large earthquake is not always the rule. Seismic quiescence may have deep implications by showing that one should not take too strictly the viewpoint that earthquakes are associated with criticality. The earthquake may not be the divergence of some accelerating events that culminate in a main shock, but instead appear after some unusually quiet period. Our model of aging shows that such a behavior may emerge from what could, at first glance, be considered as a secondary effect, the aging of the contacts, which in turn may strongly alter the distribution of the thresholds at which contacts break.

Of course one has to be careful in extending conclusions drawn from a simple, one-dimensional model to the complexity of natural earthquakes. As discussed above we hope that such a model can nevertheless give useful hints for further studies of real earthquakes. But it is also obvious that developments of the model are

clearly needed, in particular, generalizations to 2D models of the interface and 3D models of the elastic slider which would improve the GR statistics and provide a power-law for earthquake spatial distribution. The process describing the aging of the contacts also certainly needs further investigations as it is presently largely arbitrary. We use a stochastic process that goes to a Gaussian distribution in one limit and a power law in another, but how this process can be generated by the dynamics of “macro-contacts” is still open.

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APPENDIX A: METHODS

A1 Parameters

We use dimensionless units, where four model parameters are set to 1, $k = 1$, $\lambda_c = 1$, $t_0 = 1$ and $f_s = 1$ so that the other parameters should be compared with those. The IL consists of $N = 101$ λ_c^3 -cubes, each of mass $m = 10^{-4}$, as chosen in our analysis of stick-slip phenomena (Braun, 2015). The springs between the cubes have the elastic constant $K = 100$. The UL consists of parallelepipeds of height $N_L = 25$ (so that their masses are $N_L m$); for the Poisson ratio we took $\sigma_P = 0.3$. With these parameters, the crack propagation distance (where aftershocks are expected) is $\Lambda \sim 30$ (Braun & Scheibert, 2014).

For the rate of aging we took $\beta = 100$, and for the parameter ε we used $\varepsilon = 2/\Delta f_s^2$. With these parameters the statistics of the magnitudes of earthquakes follow the GR law for an interval of magnitudes $\Delta \mathcal{M} \sim 2$, which is of course limited by the size of the model.

For the dispersion of thresholds we took $\Delta f_s = 0.3$, and used $\nu = 2$ to mimic a delay in contact formation.

The initial stretching of the contacts as well as the initial thresholds of the newborn contacts, when they form again after breaking, are taken from a Gaussian distribution with zero mean and dispersion Δf_s . The UL is moved through the springs with elastic constant $K_d = 0.5$ attached to its top and driven with the velocity $v = 0.1$. The equations of motion are solved by a Runge-Kutta method with a viscous damping $\eta = 0.3\omega_0$, where $\omega_0 = (K/m)^{1/2}$, so that the dynamics of the system is underdamped. With these parameters, the motion of a single λ_c^3 -block exhibits a stick-slip behavior (recall that stick-slip exists only for a certain interval of driving velocities (Braun, 2015)).

The duration of the runs is $t_{\max} = 2 \times 10^6$. The first 33% of a trajectory are discarded to allow the system to lose memory from the initial configuration and reach a steady state (dashed vertical line in Fig. 3).

A2 Protocol and data analysis

. In a simulation we calculate and store the driving force $F(t)$ and the dimensionless quake amplitude defined as the sum of the force drops due to broken contacts at every time step,

$$\mathcal{A}(t) = \sum_{i=1}^N \Delta F_i(t) / (N f_s). \quad (\text{A.1})$$

The rate of shocks, $n(t)$, is defined as the number of shocks per time unit regardless of their amplitude. The dimensionless earthquake magnitude is defined as

$$\mathcal{M} = \frac{2}{3} \log_{10} (\mathcal{A} / \mathcal{A}_0), \quad (\text{A.2})$$

where \mathcal{A}_0 is a constant which defines the magnitude scale.

As for the study of real earthquakes (Bouchon et al, 2013) the analysis of foreshocks is delicate because one has to distinguish “main shocks” in the middle of many events, occurring randomly and with a broad range of amplitudes. To be qualified as “foreshocks” events must be related to the main shock both in time and space, and moreover a large foreshock should not be considered as an independent MEQ. In analyzing the results, we used the following protocol (Braun & Tosatti, 2014). First we remove small “background” earthquakes with amplitudes below some level from the analysis. We only retain $\mathcal{A} > 2 \langle \mathcal{A}(t) \rangle$ (see broken red line in Fig. 3). Next, we single out the main shocks above some level \mathcal{A}_{MEQ} . We took here $\mathcal{A}_{\text{MEQ}} = \kappa \mathcal{A}_{\text{max}}$, with $\kappa = 0.03$ (solid red line in Fig. 3). We also used the following rescaling procedure to detect large events which could be related. If n_{MEQ} is the total number of MEQs in a simulation, let us call $\sigma = S / n_{\text{MEQ}}$ the average area occupied by a single MEQ in the (t, x) plane, with $S = N \lambda_c t_{\text{max}}$. We then rescale the time coordinate $t \rightarrow \tau = t / \alpha$ with $\alpha = \sigma / \lambda^2$, where $\lambda \sim \Lambda$ is some distance chosen in such a way that the distribution of MEQs on the (τ, x) plane becomes *isotropic* ($\lambda \approx 30$ and $\alpha \approx 300$ for the parameters used in Fig. 3). In this rescaled space, events which are related to a particular MEQ lie within a circle around this MEQ so that they can be easily identified. We can now scan all MEQ coordinates on the (τ, x) plane and, if the distance $\rho_{ij} = [(\tau_i - \tau_j)^2 + (x_i - x_j)^2]^{1/2}$ between two MEQs i and j is smaller than some value ρ_{cut} then only the larger of these two MEQs remains as the MEQ, while the lower one is removed from the list of MEQs. We chose $\rho_{\text{cut}} = 0.4\lambda \sim 0.4\Lambda$, i.e. to be considered as related to each other two events must not be separated by more than about $0.4 \times$ the average distance along which a crack propagates.

With this protocol we have obtained a set of well separated MEQs isotropically occupying the (τ, x) plane, and we may calculate the temporal and spatial distribution of all earthquakes within some area around each MEQ. We consider that all the events separated from the corresponding MEQ by less than $\rho_0 = \rho_{\text{cut}} / 3$ are related to this particular MEQ. They can be either foreshocks or aftershocks. However we have to make sure that some events that occur before one of the MEQ do not belong to the tail of aftershocks following a violent MEQ which occurred earlier. This is done by examining only MEQs which are separated by a time interval long enough to allow the tail of aftershocks to die out. In practice only MEQs separated by a τ interval greater than $0.4 \times$ the full τ interval that we scan around MEQs are analyzed. This lead us to discard more than 30% of the events that would otherwise qualify as MEQs.

Then we collapse all data together, designating $\tau = 0$ for every main shock and normalizing shocks amplitudes on the corre-

sponding main shock value. We compute $n(\tau)$ the number of secondary events (foreshocks or aftershocks) in a given $\delta\tau = 0.0648$ interval averaged over all the investigated MEQs as well as the standard deviations of their values for the different MEQs, which provides an estimate of the error bars for $n(\tau)$ and gives us an information on the fluctuations of the earthquake activity in a given τ interval around a MEQ.